

# Overall Outline

- **Lecture I: Observations and planetary flow theory (GFD<sup>(⌘)</sup>)**
- ➔ **Lecture II: Atmospheric LFV<sup>(\*)</sup> & LRF<sup>(\*\*)</sup>**
- **Lecture III: EBMs<sup>(+)</sup>, paleoclimate & “tipping points”**
- **Lecture IV: The wind-driven ocean circulation**
- **Lecture V: Advanced spectral methods—SSA<sup>(±)</sup> *et al.***
- **Lecture VI: Nonlinear & stochastic models—RDS<sup>(◇)</sup>**

(⌘) GFD = Geophysical fluid dynamics

(\*) LFV = Low-frequency variability

(\*\*) LRF = Long-range forecasting

(+) EBM = Energy balance model

(±) SSA = Singular-spectrum analysis

(◇) RDS = Random dynamical system

# **Lecture II: Atmospheric Low-Frequency Variability (LFV) & Long-Range Forecasting (LRF)**

## **Outline**

1. Observations of **persistent anomalies**
  - **Blocked** & **zonal** flows
  - Characteristics of **persistent anomalies**
2. The **deterministic chaos** paradigm
  - **Forced** dissipative systems
  - **Successive bifurcations**
  - **Predictability** and **prediction**
3. “**Waves**” vs. “**particles**”
  - **Multiple regimes** & Markov chains
  - **Oscillatory modes** & broad spectral peaks
  - Which one is it & **how does that help?**

## ***Lecture II: Outline***

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- Characteristics of persistent anomalies

### 2. The deterministic chaos paradigm

- Forced dissipative systems
- Successive bifurcations
- Predictability and prediction

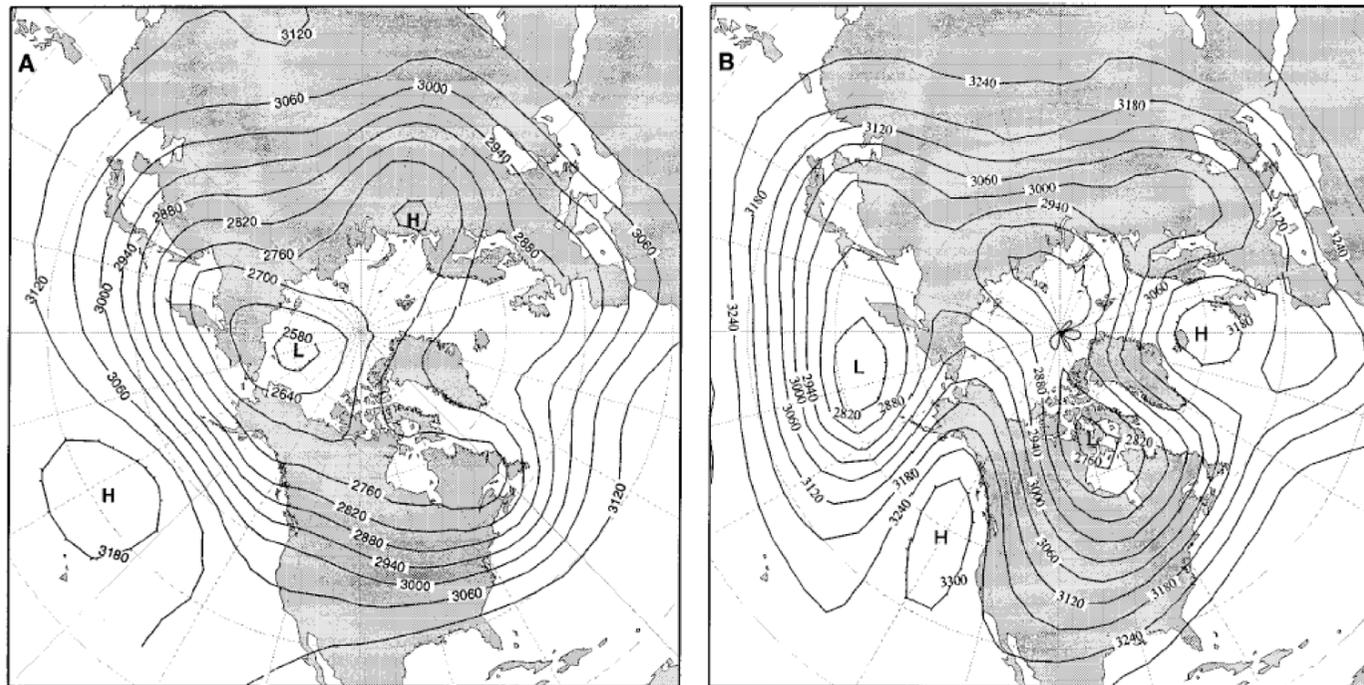
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# Transitions Between Blocked and Zonal Flows in a Barotropic Rotating Annulus with Topography

**Zonal Flow**  
13–22 Dec. 1978

**Blocked Flow**  
10–19 Jan. 1963



**Fig. 1.** Atmospheric pictures of (A) zonal and (B) blocked flow, showing contour plots of the height (m) of the 700-hPa (700 mbar) surface, with a contour interval of 60 m for both panels. The plots were obtained by averaging 10 days of twice-daily data for (A) 13 to 22 December 1978 and (B) 10 to 19 January 1963; the data are from the National Oceanic and Atmospheric

Administration's Climate Analysis Center. The nearly zonal flow of (A) includes quasi-stationary, small-amplitude waves (32). Blocked flow advects cold Arctic air southward over eastern North America or Europe, while decreasing precipitation in the continent's western part (26).

Weeks, Tian, Urbach, Ide, Swinney, & Ghil (*Science*, 1997)

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# **Characteristics of intraseasonal variability**

## **(~ atmospheric LFV)**

- 1. Geographically fixed appearance and regional character (\*)**  
(“teleconnections” – Wallace & Gutzler, 1981)
- 2. Persistence**  
(*persistent anomalies* – Dole, 1982, 1986; Horel, 1985)
- 3. Recurrence**  
(*multiple regimes* – Mo & Ghil, 1987, 1988; Kimoto & Ghil, 1993a,b)
- 4. Barotropic structure**  
( barotropic, or 3<sup>rd</sup>, adjustment; see next page)

(\*) but Branstator (1987) & Kushnir (1987), 25-day hemispheric wave;  
Benzi et al., 1984 +, hemispheric bimodality;  
Wallace, Thompson & co. – Arctic Oscillation.

# ***Barotropization***

– barotropic (3rd) adjustment<sup>(\*)</sup>

**(a) statistical theory of turbulence**

(Charney, 1971; Rhines, 1979; Salmon, 1980)

**(b) evolution of baroclinic eddies & "wave maker"**

(Hoskins & Simmons, 1978; Green-Ilari-Shutts)

**(c) external Rossby wave, & its instability**

(Held-Panetta-Pierrehumbert, 1985–87)

<sup>(\*)</sup>After hydrostatic (1st) and baroclinic (2nd) adjustment.

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# Forced dissipative systems

Most fluid dynamical problems — and many other problems in biology, chemistry, and continuum physics — lead to ODEs (or equivalent PDEs) of the form

$$\dot{x}_i = a_{ijk}x_jx_k - b_{ij}x_j + c_i, \quad i = 1, 2, \dots, N. \quad (\text{FD})$$

Here we used the summation convention for repeated indices. In fluid-flow problems, the quadratic terms in (FD) above represent the nonlinear advection term  $\vec{u} \cdot \nabla \vec{u}$ . This term is associated with the Jacobian in the QG equation.

The above equation is *autonomous* and it has unique solutions for all initial data (ID)  $x(0) = x_0$ ; these solutions depend continuously on the ID,  $x = x(t; x_0)$ . When the solutions exist for all times,  $-\infty < t < \infty$  (\*), then Eqs. (FD) define a *differentiable dynamical system* (DDS). In particular, we shall assume that this system is *forced*,  $c_i \neq 0$ , and *dissipative*,  $b_{ij}x_i x_j > 0$ .

N.B. The quadratic terms are necessarily *energy conserving* if  $a_{ijk} = -a_{ikj}$ . and the orbits of (FD) describe a flow in the phase space of  $\{x_i, i = 1, \dots, N\}$ .

(\*) *Counterexample*. The solutions of  $\dot{x} = x^2$  are unique and depend continuously on  $x_0$  but they blow up at  $t = 1$ !

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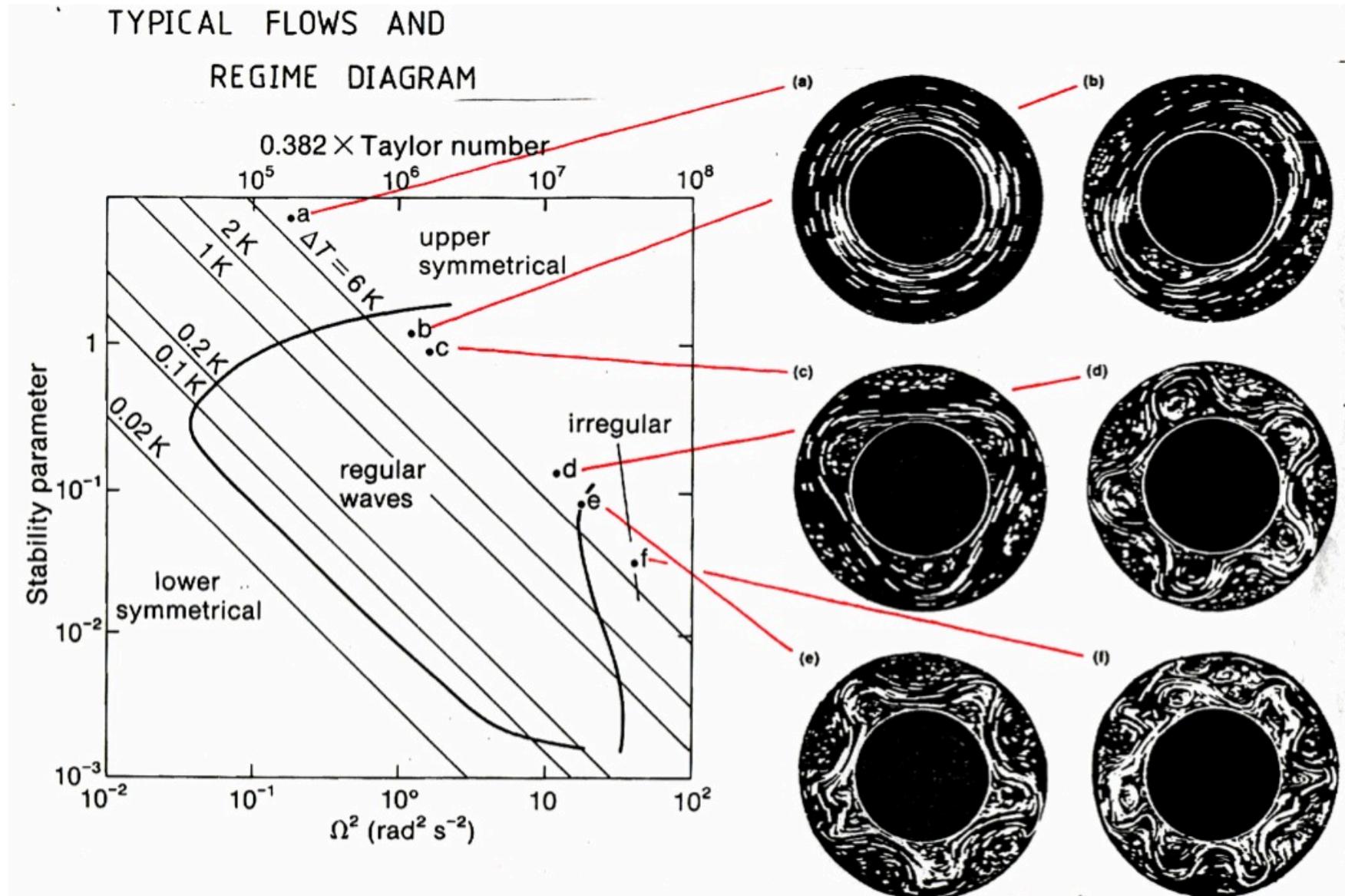
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# Rotating Convection: An Illustration



# GFD, bifurcations and chaos

**Problem 3:** Read the paper listed below and report to the class on its contents.

Ghil, M., P. L. Read and L. A. Smith, 2010: Geophysical flows as dynamical systems: The influence of Hide's experiments, *Astron. Geophys.*, **51**(4), 4.28–4.35

## General idea

As we push the system harder, it responds by coming up with more complex responses, i.e., **it loses symmetry** in both time & space. **In time**, it may go from being in steady state to being periodic and then chaotic; **in space**, it often goes from being homogeneous to periodic and then to irregular. thus, the two kinds of symmetry loss are interrelated.

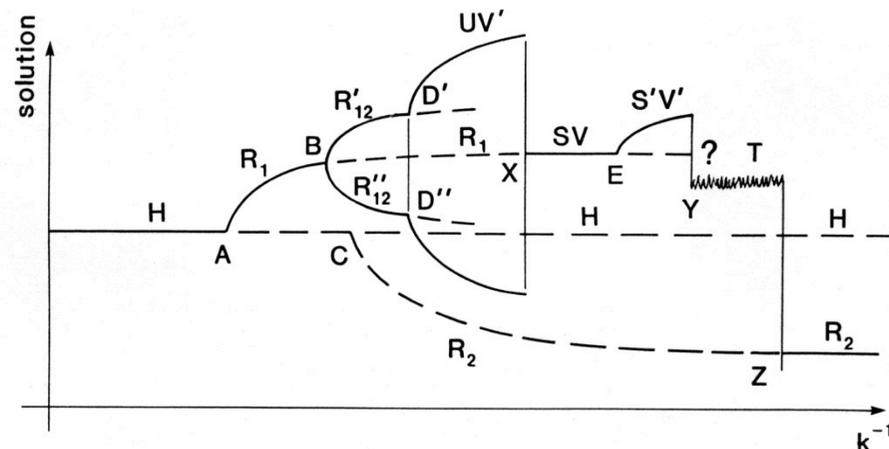
## Bifurcation diagram

### General situation

$$u_t = N(u; \mu)$$

$$N(u_0; \mu_0) = 0.$$

- 1) If  $L_0 = N/\partial u$  at  $(u_0; \mu_0)$  is nonsingular, then a unique **branch** of solutions  $u = u(\mu)$  through it exists and is given by  $u \cong u_0 + (\partial u/\partial \mu)|_{u=u_0} \mu - \mu_0$ .



- 2) The points at which  $\det L_0 = 0$  (i.e., where the Implicit Function Theorem fails) are called **bifurcation** points, and they are in general **isolated**. Near such points, the behavior of (2 or more) solutions is parabolic:

$$u - u_0 \sim (\mu - \mu_0)^{1/2}$$

***Calm in the face of chaos ...***

46

***But just wait till we bring  
in randomness, too!***



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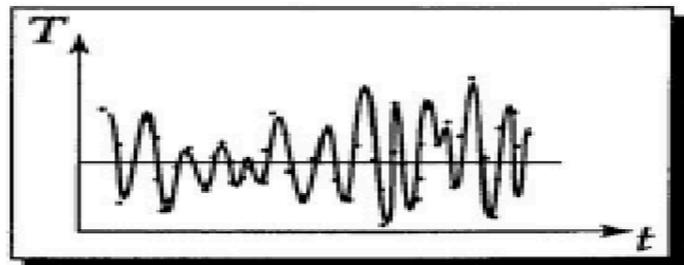
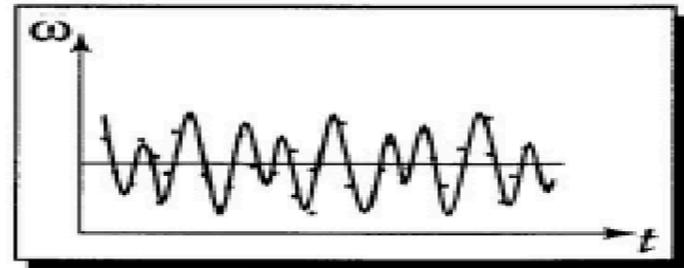
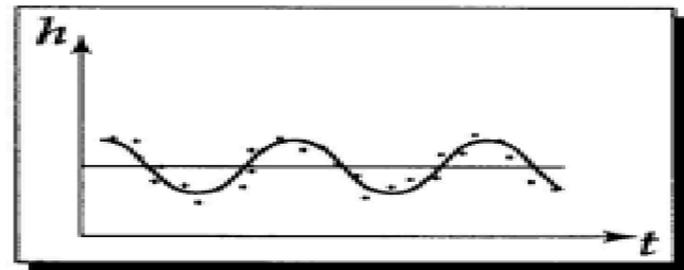
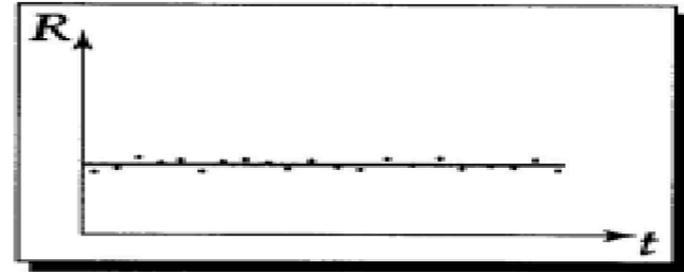
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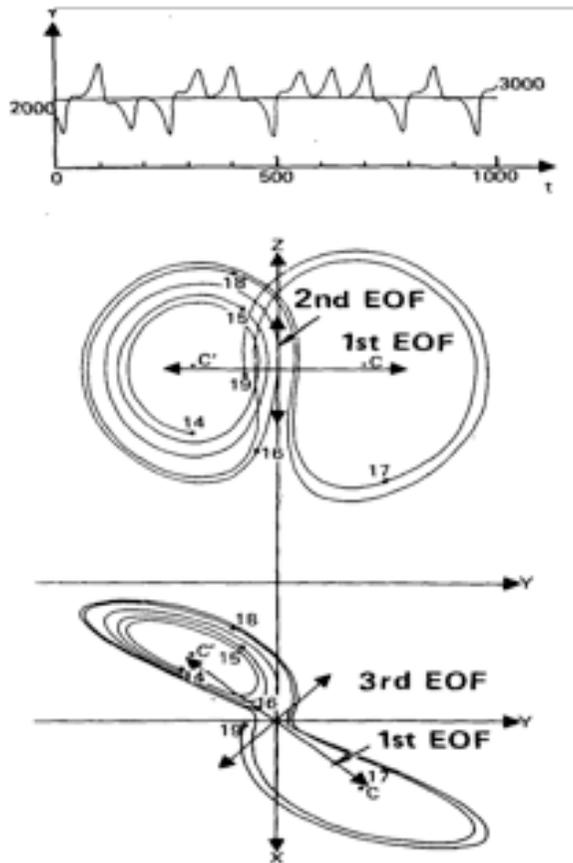
## Predictability & prediction

1. Easiest to predict: a constant state, e.g., Earth's radius  $R$  → one needs only one number.
2. A little harder: periodic phenomena, e.g., sunrise, the tides → this requires 3 numbers — the period  $p$ , the amplitude  $A$  & the phase  $\phi$ , in this order.
3. Even harder: quasi-periodic phenomena, e.g. the planetary orbits in celestial mechanics → we need  $3n$  numbers, where  $n$  is the numbers of periodic orbits, which may be large but finite.
4. And so how about some real stuff, like thermal convection, weather, the markets → one needs an infinity of numbers.

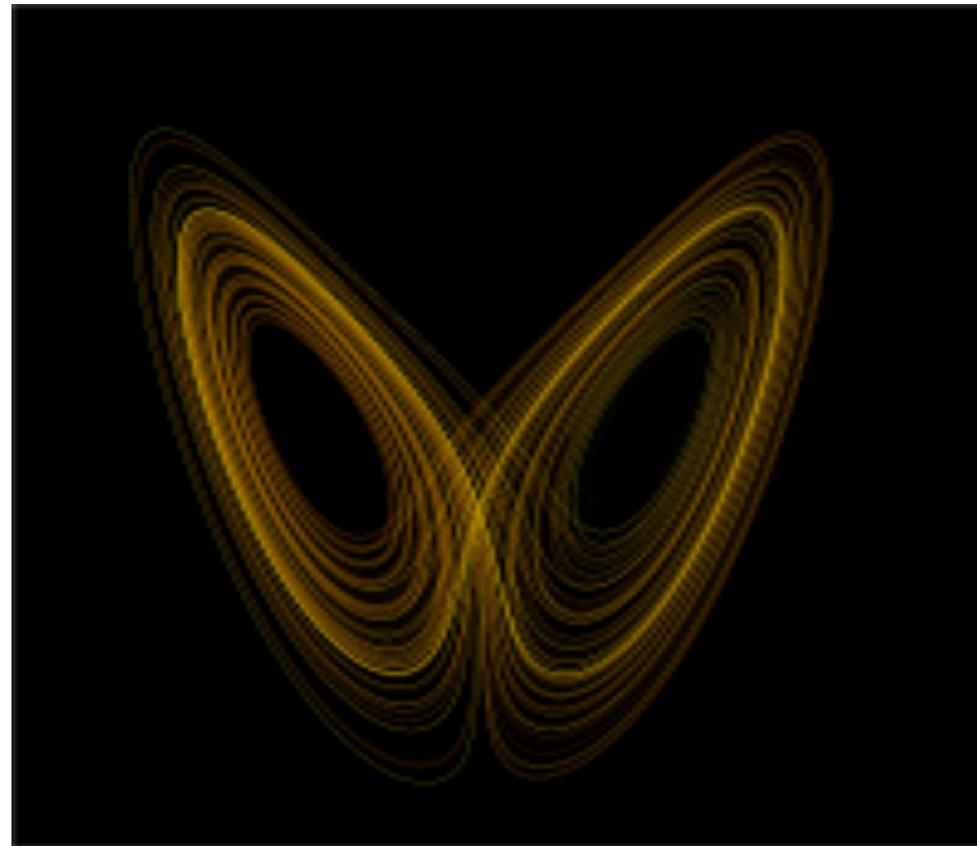
🍏 The more complex the phenomenon, the harder it is to predict.



# The Lorenz convection (1963a) model – some numerical solutions



Plot of  $Y = Y(t)$  + projections  
onto the  $(X, Y)$  &  $(Y, Z)$  planes



Trajectory in phase space

Both for the canonical “chaotic” values  $\rho = 28$ ,  $\sigma = 10$ ,  $\beta = 8/3$ .

# The Lorenz (1963a) convection model

**Problem 4:** Find the appropriate software to compute the statistics of the Lorenz “butterfly” – e.g., pdf, EOFs – and use it to do so.

## *Glossary*

pdf = probability density function

EOF = empirical orthogonal function

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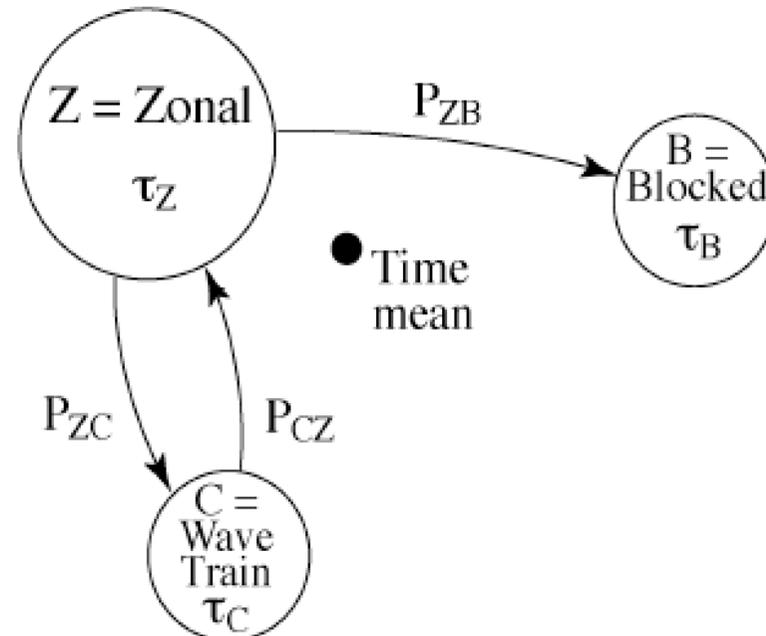
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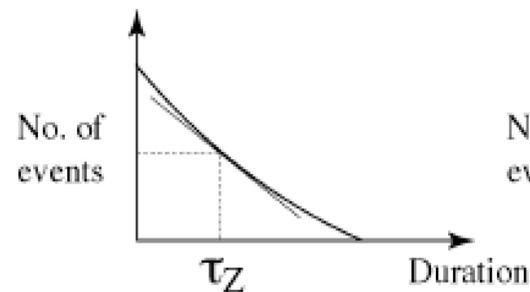
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# Coarse-graining *Markov-chain description of LFV*

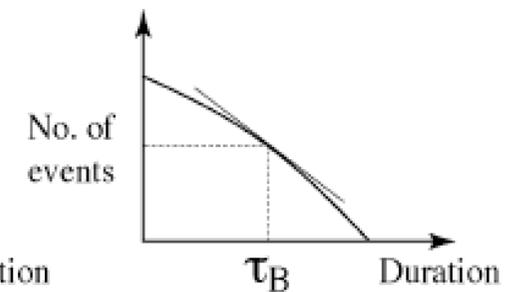
1. Reduce the number of degrees of freedom to the most important ones – highest variance.
2. Describe the dynamics in this reduced subspace.



(a)



(b)



(c)

# Multiple Flow Regimes

## A. Classification schemes

### 1) By position

#### (i) *Cluster analysis*

##### – categorical

– NH, Mo & Ghil (1988, *JGR*) – fuzzy

– NH + sectorial, Michelangeli et al. (1995, *JAS*) – hard (*K*-means)

##### – hierarchical

– NH + sectorial, Cheng & Wallace (1993, *JAS*)

#### (ii) *PDF estimation*

##### – univariate

– NH, Benzi *et al.* (1986, *QJRMS*); Hansen & Sutera (1995, *JAS*)

##### – multivariate

– NH, Molteni *et al.* (1990, *QJRMS*); Kimoto & Ghil (1993a, *JAS*)

– NH + sectorial, Kimoto & Ghil (1993b, *JAS*);

Smyth *et al.* (1999, *JAS*)

After Ghil & Robertson (2002, *PNAS*)

# Multiple Flow Regimes (continued)

## A. Classification schemes (continued)

### 2) By persistence

#### (iii) *Pattern correlations*

- NH, Horel (1985, *MWR*)
- SH, Mo & Ghil (1987, *JAS*)

#### (iv) *Minima of tendencies*

- Models: Legras & Ghil (1985, *JAS*); Mukougawa (1988, *JAS*);  
Vautard & Legras (1988, *JAS*)
- Atlantic- European sector : Vautard (1990, *MWR*)

## B. Transition probabilities

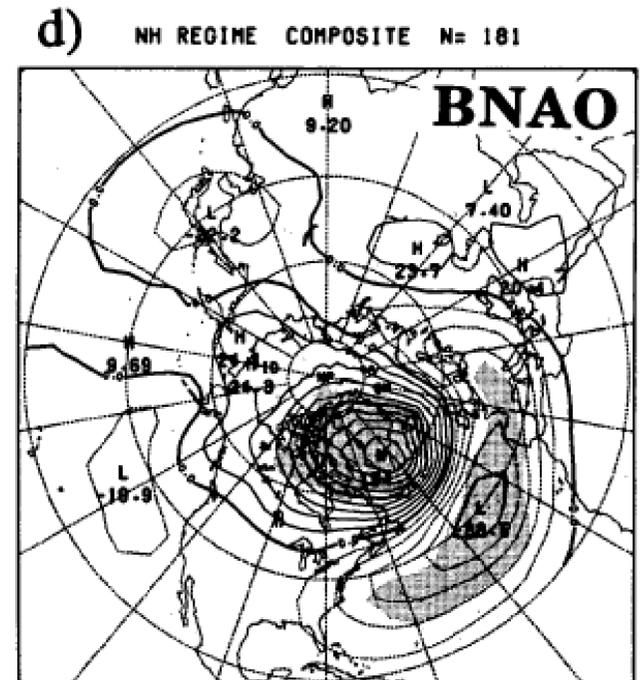
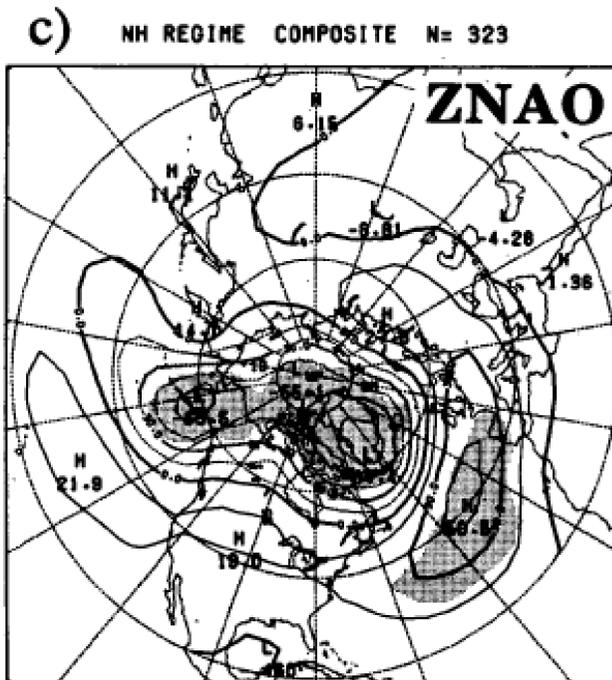
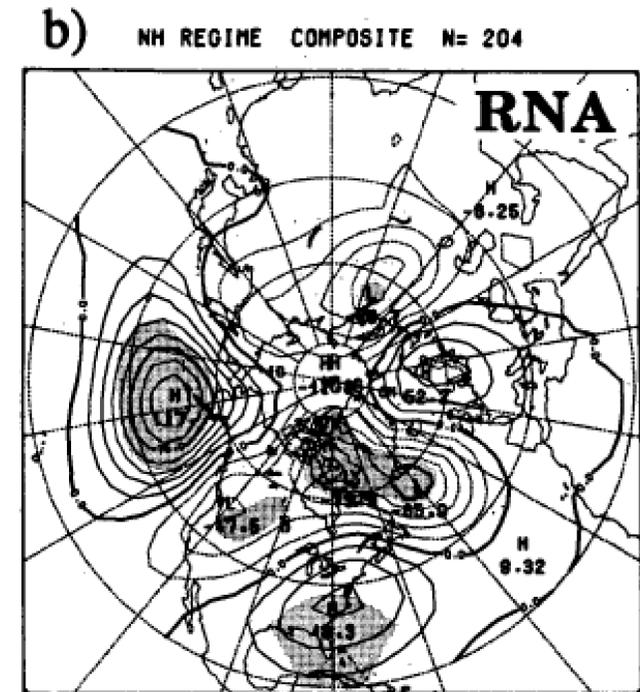
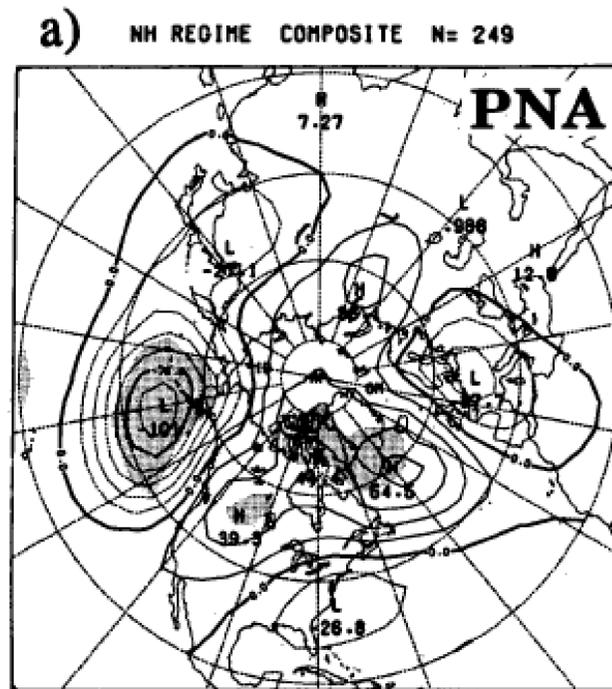
- (i) Model & NH – counts (Mo & Ghil, 1988, *JGR*)
- (ii) NH & SH – Monte Carlo (Vautard *et al.*, 1990, *JAS*)

After Ghil & Robertson (2002, *PNAS*)

## Multiple Flow Regimes

– lowest common denominator, I

Four regimes:  
blocked vs. zonal,  
in the Pacific–North-  
American (PNA) & the  
Atlantic–European  
sector, respectively  
(Kimoto & Ghil,  
*JAS*, 1993a)



## Multiple Flow Regimes

– lowest common denominator, II

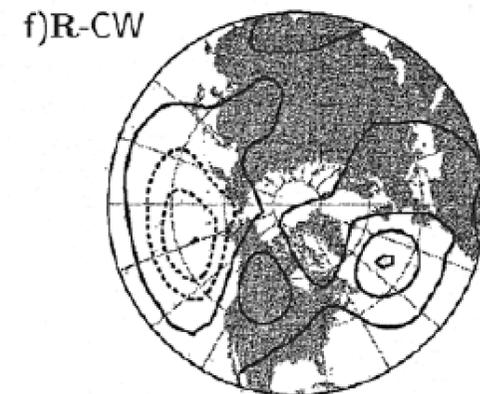
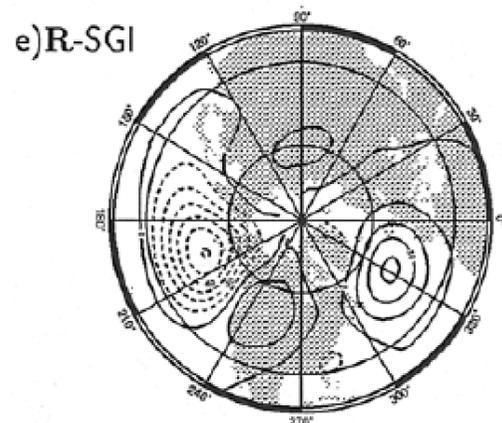
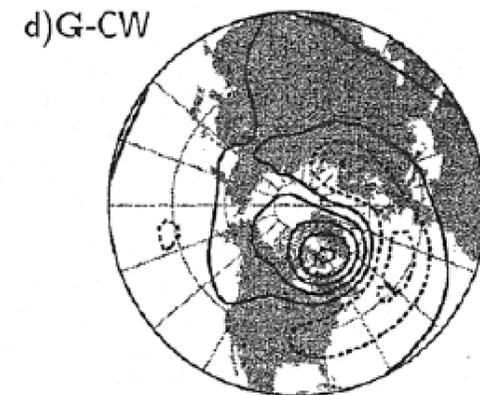
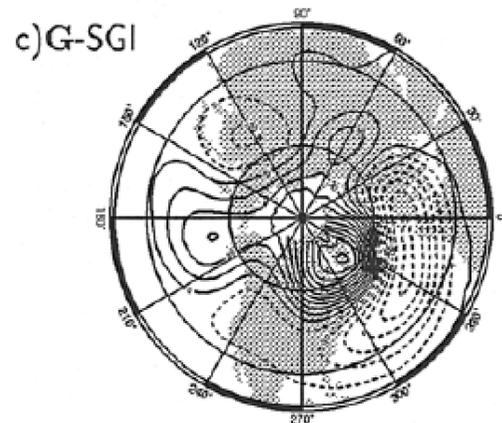
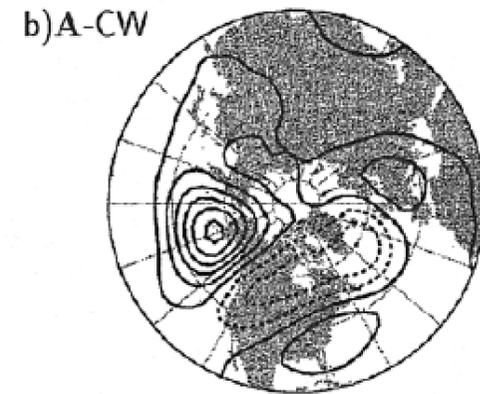
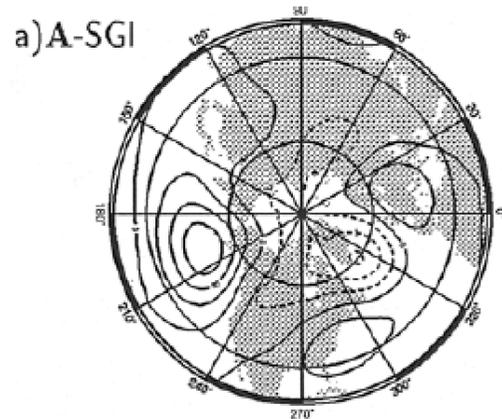
Cheng & Wallace  
(*JAS*, 1993; **CW**) &  
Smyth, Ghil & Ide  
(*JAS*, 1997; **SGI**) agree  
well on 3 of the 4 regimes  
in Kimoto & Ghil  
(*JAS*, 1993a; **KG**):

**A** – Gulf of Alaska ridge ~  
KG's **RNA**

**G** – high over Greenland ~  
KG's **PNA**

**R** – enhanced ridge over  
Rockies ~ **BNAO**

SGI's sectorial analyses  
also find KG's **ZNAO** to  
be quite robust.



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# Waves vs. Particles: A Pathway to Prediction?

Is predicting as hard  
as it is claimed to be?

No, it's actually quite easy:

Just flip a coin or roll a die!

What's difficult, though, is

trusting the prediction

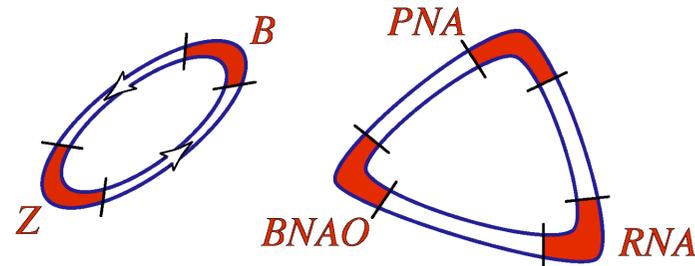


*That's where a little  
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Based on Ghil & Robertson (2002)

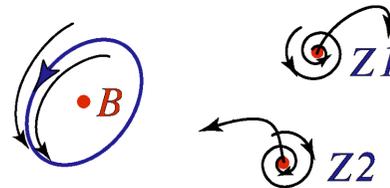
## "Waves vs. Particles" in Atmospheric Low-Frequency Variability

1. Are the regimes but slow phases of the oscillations?



Kimoto & Ghil  
(JAS, 1993a, b)

2. Are the oscillations but instabilities of particular equilibria?

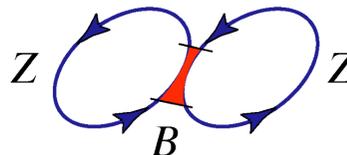


Legras & Ghil  
(JAS, 1985)

3. How about both: "chaotic itinerancy" (Itoh & Kimoto, JAS, 1999)

4. How about **neither?** Null hypotheses:

a) It's all due to interference of linear waves, e.g.,  
neutrally stable Rossby waves;



Lindzen *et al.*  
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b) It's all due to red noise — Hasselmann (*Tellus*, 1976),  
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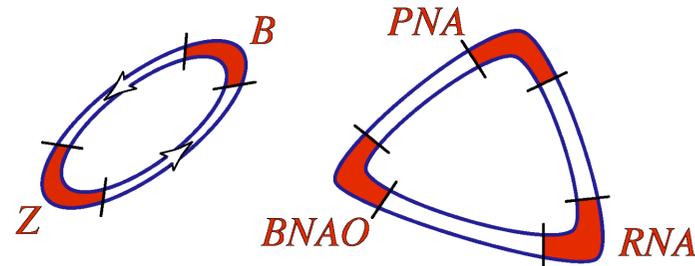


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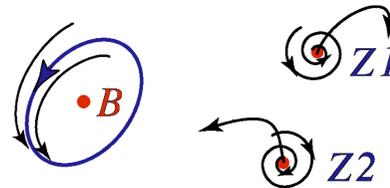
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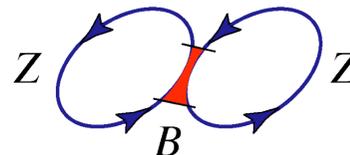


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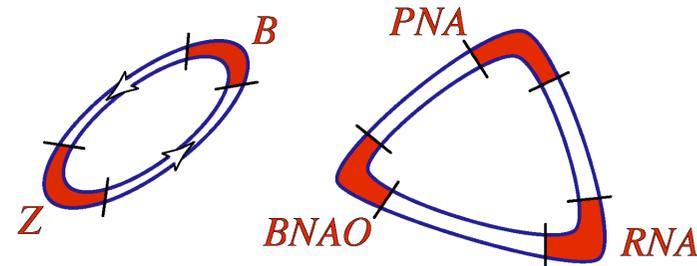
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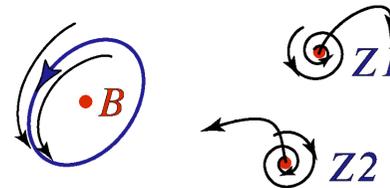
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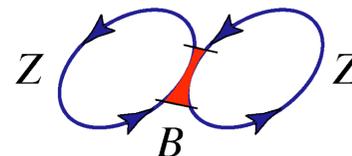


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**Reserve slides**

## ***Lecture I: Outline***

1. General introduction and bibliography
  - Scale dependence of atmospheric & oceanic flows
  - Turbulence & predictability
2. Basic facts of large-scale atmospheric life
  - The atmospheric heat engine
  - Shallowness
  - Rotation
3. Flow regimes, bifurcations & symmetry breaking
  - The rotating, differentially heated annulus
  - Regime diagram & transitions

## ***Lecture II: Outline***

### 1. Observations of **persistent anomalies**

- **Blocked** & **zonal** flows
- Characteristics of **persistent anomalies**

### 2. The **deterministic chaos** paradigm

- **Forced** dissipative systems
- Successive bifurcations
- **Predictability** and **prediction**

### 3. “**Waves**” vs. “**particles**”

- **Multiple regimes** & Markov chains
- **Oscillatory modes** & broad spectral peaks
- Which is one is it & **how does that help?**